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# DELIVERABLE

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#### 23 1 Abstract

To better understand the effect and influence of spatial density of ground motion 24 predictions, we attempted to combine data from multiple seismic networks from the same 25 region and compare intensity predictions with and without new seismological networks. The 26 U-Net neural network architecture is used as a ground motion model that predicts the mean 27 and standard deviation of a target intensity measure (IM, PGA) in the form of maps. The 28 U-Net can interpolate the intensity measures between the observation points inherently and 29 we here try to analyse the effect of spatial density in this interpolation by using multiple 30 networks. Kanto basin in Japan is selected for this study due to the availability of dense 31 seismic networks within the basin and we integrated the KIK-net and the MeSO-net networks 32 within the basin for this study. The results show that the errors and the uncertainty levels 33 between the predictions and the observations at the interpolated sites slightly decreased 34 after adding an additional network for the training process, but the statistical significance is 35 questionable. However, more networks or more observations per station are required to be 36 integrated and trained within the same region to analyse and validate our results. 37

## 38 2 Introduction

Over the last few decades, a large number of ground-motion models (GMMs) have 39 been developed for use in seismic hazard and risk applications all over the world. GMMs 40 define the ground motion field in terms of earthquake source characteristics (e.g., magnitude, 41 fault mechanism), wave propagation (e.g., epicentral distance), and site effects (e.g., site class 42 or VS30) to predict specific intensity measures (IM). For various tectonic regions, tens or 43 even hundreds of candidate models are available due to the intricacy of the earthquake 44 process, wave propagation, and site effects; examples can be found in Douglas (2020). These 45 models represent the distribution of ground-motion in terms median intensity measure and 46 its standard deviation (Strasser et al. 2009). 47

The U-Net neural network architecture has been used recently as a ground motion 48 model that predicts the mean and standard deviation of a target intensity measure (IM, 49 PGA) in the form of maps (Lilienkamp et al. 2022). The U-Net, it learns the relationship 50 between predictive parameters and target parameters using a large number of instances 51 that are presented to train the neural network, just like all supervised learning techniques. 52 In ground-motion modelling, the target parameter is a ground-motion IM, which may be 53 deduced from predictive factors such as the moment magnitude  $M_w$  and the hypocentral 54 distance  $r_{hyp}$ . The U-Net neural network architecture has been used recently as a ground 55 motion model that predicts the mean and standard deviation of a target intensity measure 56 (IM, PGA) in the form of maps (Lilienkamp et al. 2022). The U-Net can interpolate the 57 intensity measures between the observation points inherently and we here try to analyze 58 the effect of spatial density of observation points in this interpolation by using multiple 59 seismological networks (e.g. MeSO-NET). 60

To better understand the effect and influence of arrays spatial density on ground mo-61 tion predictions, we attempted to combine data from multiple seismological networks from 62 the same region and compare intensity predictions with and without new seismological sta-63 tions. The relative surface fault motion for recording stations situated on either side of a 64 causative fault, soil liquefaction, landslides, and the general transmission of the waves from 65 the source through the various earth strata to the ground surface can all contribute to the spa-66 tial variability in seismic ground motions (Zerva & Zervas 2002). This architecture's inherent 67 capacity to process data in the form of 2D arrays (maps) makes it particularly appealing for 68 ground-motion modeling, as it allows for native operations on map data, which preserves the 69 underlying spatial distribution of ground-motion observations (Lilienkamp et al. 2022). This 70 visualization method provides a clear picture of how spatial density of observation points 71 affects ground motion predictions. This spatial variation of seismic ground motions has 72 started being analyzed after the installation of these kinds of dense instrument arrays. The 73 sensors used for these arrays are costly and also time-consuming, hence studies are required 74

<sup>75</sup> to really understand the effect of these arrays and to find an optimal array that minimises <sup>76</sup> the error in model predictions. In this study, we focus on the effect of spatial density of <sup>77</sup> the stations in the intensity predictions and the associated uncertainties by integrating data <sup>78</sup> from different seismological networks.

In this report, we use the U-net architecture to understand the relation between the IM (here PGA) and predictive parameters in the Kanto basin using a subset of the Kiban–Kyoshin (KiK-net) dataset. We then look at how integrating two datasets from the same region (Kanto Basin) - the KiK-net and the highly dense Metropolitan Seismic Observation network (MeSO-net), affects the interpolation of IMs and their associated uncertainties between the observation points. Here we work with the PGA values while the original work of (Lilienkamp et al. 2022) focused and tested only for SA(T=1s).

# 86 3 Data Used

Kanto basin in Japan is selected for this specific study due to the high-density seismic networks available within the basin. The basin has KiK- et stations, MeSO-net stations (Sakai & Hirata 2009), K-net stations (National Research Institute for Earth Science and Disaster Resilience, 2019) and some QuakeSaver devices are also installed within some buildings in the region.

We are primarily interested in determining and analyzing the difference in the uncertainty in U-Net's interpolation of intensity measures between station locations when data from numerous networks is used instead of a single network, hence, the goal of our study necessitates the use of high-density networks of this type. We focus on two networks in this report, i.e, KiK-net and MeSO-net as shown in Figure 1. Both KiK-net and MeSO-net are operated by National Research Institute for Earth Science and Disaster Resilience, NIED (2019).



**Figure 1:** KiK-net and MeSO-net station coverage in the Kanto Basin. Blue triangles indicate the MeSO-net locations (highly dense) and the red circles are the KiK-net station locations.

# 99 3.1 KiK-net dataset

KiK-net consists of pairs of strong-motion seismographs installed in a borehole and 100 on the ground surface. Here, we used exactly the same dataset given by Bahrampouri et al. 101 (2021) comprising both surface and borehole sensors. We downloaded the processed data by 102 Bahrampouri et al. (2021). The database utilized in this study comprises all earthquakes 103 with a magnitude greater than three that were recorded on the KiK-net website between 104 1996 and the end of 2017 and used by Bahrampouri et al. (2021), Lilienkamp et al. (2022). 105 We used 46,191 records from 2864 events recorded at 65 different stations. The selected KiK-106 net stations have an average inter-station distance of about 94 km. As our target intensity, 107

we used the geometric mean of the PGA values. The distribution of the dataset is given in
Figure 2.

# <sup>110</sup> 3.2 MeSO-net dataset

MeSO-net is made up of around 300 observation stations in the Tokyo metropolitan 111 region, as illustrated in Figure 1 (blue triangles). These stations are made up of five dense 112 linear arrays spaced around 2 to 3 km apart and a sparser distribution with a radius of about 113 80 km spaced roughly 4 to 10 km apart. This network is highly dense compared to the inter-114 station distance of KiK-net stations. MeSO-net stations (three-component accelerometer) 115 are located at the bottom of 20 m deep boreholes. This data is converted to correspond to 116 the intensity at the ground surface according to Aoi et al. (2021) by adding 0.5 to  $\log(PGA)$ 117 at all stations at all stations (accounting for the free surface factor). 118



**Figure 2:** Data from Kanto basin used for this study. Hit counts computed for the data distribution, dividing the distance range into 20 equally spaced bins over the hypocentral distance and the moment magnitude (Mw). a) KiK-net dataset b) MeSO-net dataset.

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The MeSO-Net waveform data is downloaded from the National Research Institute for Earth Science and Disaster Resilience (NIED) website. The website provides the data from 2021 and 2022. A total of 1921 recordings are obtained from 19 events recorded at 300 different stations in the Kanto Basin. Figure 2 depicts the data distribution of both networks - the KiK-net and the MeSO-net. Data used from MeSO-net for this study is only <sup>124</sup> available from a limited time period.

# 125 4 Methodology

## 126 4.1 The U-net architecture

The U-net neural network architecture (Ronneberger et al. 2015) is the central com-127 ponent of our methodology, as it can offer us with a fully non-ergodic and data-driven GMM. 128 Lilienkamp et al. (2022) introduced this to the ground motion modelling area, and the paper 129 tests the U-net as a GMM and explains how it can work as a GMM. Figure 2 illustrates this. 130 The input parameters are the nine maps shown in the figure, i.e., the latitude and longitude 131 of the event, the depth, magnitude, the coordinates of each pixel in the input layer (572) 132  $\times$  572 pixels), the hypocentral distance and the depth to seismic bedrock. The output will 133 be the mean,  $\hat{y}$  and variance,  $\sigma^2$  of the IM. The valid convolution operations reduces the 134 resolution of the output features is to  $388 \times 388$  compared to that of the input features. 135 Through a series of instances and iterations, the U-net algorithm learns the link between the 136 input parameters and the output. Detailed information on the neural networks can be found 137 in LeCun et al. (2015). The U-net is particularly useful for ground motion modelling since 138 it is designed to process data in the form of 2D arrays, or maps. The U-Net's predictive 139 parameters are provided in the form of a stack of maps spanning a predefined area, and the 140 U-Net is trained to provide estimates of the target IM's mean and variance. The negative 141 log-likelihood is considered as the model error or the loss between the model outputs and 142 the observations for an event with N observations. The loss is iteratively minimized using 143 the gradient descent algorithm Adam (Da 2014), with the gradient effectively implemented 144 using backpropagation (Rumelhart et al. 1986). 145

After one cycle (epoch) of training using all of the training data, the neural network's training is tested using the validation dataset. After a given number of epochs, the loss on the validation dataset does not decrease any further, indicating that the training is complete and the U-Net may now be used as a GMM. Instead of point-wise observations,



**Figure 3:** This study's U-Net architecture [Figure from Lilienkamp et al. (2022)]. The U-Net receives the predictive parameters for a single earthquake as input in the form of a stack of maps. The input is processed, and the intensity measure (IM) mean  $\hat{y}$  and variance  $\hat{\sigma}^2$  estimators are provided as output. Lat<sub>e</sub>, lon<sub>e</sub>, and d<sub>hyp</sub> are the latitude, longitude, and depth of the event hypocenter, respectively. Mw is the moment magnitude, xs and ys are the coordinates of each pixel in the input layer, rhyp is the hypocentral distance, and  $z_b$ edrock is the depth to seismic bedrock. This figure is described in more detail in (Lilienkamp et al. 2022). This figure is based on the first figure of Ronneberger et al. (2015).

the U-net predictions are continuous maps. This means that the loss function is generated first at sites where actual observations are available, and then the U-Net interpolates the learned attenuation relation from those locations. The data was separated into training and validation events using the same training technique as (Lilienkamp et al. 2022), i.e.,

• The data were divided into training and validation events. All events occurring up to 2015 are used for training, while those occurring 2015 and onwards are considered validation events. • Stations are divided into a number of chunks at random  $(N_{chu})$  - for each station chunk, one U-net is trained (the selected station chunk is excluded from training in order to be used for validation after training).

• The U-net is run  $N_{init}$  times to quantify the variability caused by random coefficient initialization. The final predictions for the mean and variance of the target IM for the event e ( $\hat{Y}_e$  and the  $\hat{\Sigma}_e$ , respectively), are obtained by ensemble averaging the mean and variance predictions of the individual U-Nets ( $\hat{y}_e^{ij}$  and the  $\hat{\sigma}_e^{ij}$ , respectively). The total number of U-Nets will be,  $N_U = N_{chu} * N_{init}$  and the mean and variance of the final predicted IM according to the law of total expectation and the law of total variance (Blitzstein & Hwang 2015) for an event will be

$$\hat{Y}_e = 1/N_U \sum_{i=1}^{N_{chu}} \sum_{j=1}^{N_{init}} \hat{y}_e^{ij}$$
(1)

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$$\hat{\Sigma}_{e}^{2} = 1/N_{U} \sum_{i=1}^{N_{chu}} \sum_{j=1}^{N_{init}} [\hat{\sigma}_{e}^{ij^{2}} + \hat{y}_{e}^{ij^{2}}] - \hat{Y}_{e}^{2}$$
(2)

The U-Net learns the site amplification as a function of the coordinates of the station location, source-location specific variability from the event latitude and longitude, and pathspecific amplification from the coordinates of each pixel and from the latitude and longitude of the events. We initially analyzed the U-net findings using only the KiK-net dataset, then trained the U-net using both the KiK-net and MeSO-net datasets to see how the results differ in terms of increasing the number of observation sites between the existing ones (the KiK-net stations).

#### 175 4.2 U-Net training with different seismological networks - Kanto Basin

We started (the first stage) of the U-Net training with the KiK-net strong-motion dataset by (Bahrampouri et al. 2021) for peak ground acceleration (PGA) measurements. All events occurring up to 2015 are used for training, while those occurring 2015 and onwards are

considered validation events. The 65 stations are separated into 5 chunks, with each chunk 179 receiving a random selection of the stations and 10 random initializations. We proceed with 180 the inclusion of the MeSO-net dataset after obtaining the predicted PGA values with the 181 KiK-net dataset. We use both the KiK-net and MeSO-net recordings for the training in 182 this stage. For the training phase, both the events occurring before 2015 (only KiK-net) 183 and after 2020 (only MeSO-net events) are evaluated. The validation events are set as the 184 same sets we used in the first step in order to understand the changes after adding a new 185 dense network (MeSO-net) with the KiK-net data in training process, i.e., the same KIK-net 186 station chunks are used for validation of interpolated IM values. A summary of this is shown 187 in the flowchart. (Figure 4) 188

### 189 5 Results and Discussion

## <sup>190</sup> 5.1 Interpolation Quality

We first analyse the effect in the interpolation of IMs when arrays of stations (MeSO-191 net) are integrated to the KiK-net dataset in the same region, Kanto basin. As we have 192 discussed earlier the U-net automatically interpolates the learned relation from the obser-193 vation locations across the output area and we study the quality of this interpolation using 194 the partial ensemble estimators,  $\hat{\mathcal{Y}}^i$ , for which the *i*th station chunk was not used during 195 training. Averaging over the subsets of U-Nets that share the same ith station validation 196 chunk yields partial ensemble estimators. Lilienkamp et al. (2022) evaluated and classified 197 these partial ensemble estimators in four different ways to understand the performance of 198 U-Net at the interpolated sites and we use the same technique to assess the effect of spatial 199 density in the interpolation. The four different categories are (1) training events recorded 200 on training stations, (2) validation events recorded on training stations, (3) training events 201 recorded on validation stations, and (4) validation events recorded on validation stations. 202 We average the root mean square error (rmse) between the observations and predictions 203 (Ln) over the five station chunks  $\hat{\mathcal{Y}}^i$ . The average of the rmse values using only the KiK-net 204



Figure 4: Flowchart describing the process used in this study.

stations is given in Table 1. The rmse values are larger when using the validation stations, as seen in the table, because these predictions from the validation stations have additional errors owing to interpolation. The training events and the training stations will have the smallest rmse values as those observations are the ones used for training.

**Table 1:** Average Root Mean Square Errors  $\pm$  1 Standard Deviation between Observations and Predictions of the Five Partial Ensemble Estimators using only KiK-net.

Configuration	Rmse
Training stations/training events	$0.432\pm0.007$
Training stations/Validation events	$0.496 \pm 0.005$
Validation stations/training events	$0.871 \pm 0.104$
Validation stations/validation events	$0.854 \pm 0.090$

**Table 2:** Average Root Mean Square Errors  $\pm 1$  Standard Deviation between Observations and Predictions of the Five Partial Ensemble Estimators using both KiKnet and MeSO-net.

Configuration	Rmse
Training stations/training events	$0.429\pm0.007$
Training stations/Validation events	$0.494 \pm 0.005$
Validation stations/training events	$0.865 \pm 0.086$
Validation stations/validation events	$0.841 \pm 0.080$

The rmses in Ln units are then analyzed using both the KiK-net and MeSO-net 209 stations to see how integrating more stations affects the results. The validation stations 210 rmses with and without the MeSO-net stations are of particular interest because they are the 211 predictions with the interpolation error. The average of the rmses between the observations 212 and predictions over the five  $\hat{\mathcal{Y}}^i$  using both the networks is given in Table 2. As shown in 213 the table the rmse values decreased in all the four configurations. The change of rmses are 214 within the standard deviation provided in the tables for all the four combinations, which 215 may provide the idea that the changes are not that significant. Due to the small number of 216 MeSO-net records compared to KIK-net records, the influence on training the neural network 217 is relatively small. Thus, while the number of records is small, the MeSO-net network is 218 actually predestined to improve the quality of interpolated values in this region. We may 219

observe more significant changes in the validation stations on validation events combination
by using a more temporal dense network, eg.the Kyoshin Network (K-net), (NIED 2019).

We also try to understand the effect on the uncertainty in prediction when another 222 seismological networks are integrated to the training dataset. Lilienkamp et al. (2022) used 223  $\hat{\Sigma}$  (given in equation 2) as a proxy for the scatter in observations at station locations, as 224 it is what we learn with the neural network, and it is supposed to have learned the scatter 225 in observations and verified its reliability by comparing the distribution of the standardized 226 residuals,  $\tilde{\Delta}$ , with the standard normal distribution. The residuals are standardized by 227 dividing each individual residual by its predicted standard deviation,  $\hat{\Sigma}$ . We calculated the 228  $\Delta$  achieved using only the KiK-net data and also together with the MeSO-net dataset for 229 training. Once this is achieved we compare it with the target normal distribution to analyse 230 the difference in both distributions (Figure 5). We get a similar standard deviation of  $\Delta$ 231 in both cases, i.e around 14-15 % decrease than the targeted value. Since, we standardized



**Figure 5:** Comparison of the distribution of standardized residuals with the targeted standard normal distribution. Solid and dashed vertical lines indicate the empirical mean and standard deviation (std) of standardized residual,  $\tilde{\Delta}$  respectively. a) Training with only KiK-net b) Training with both KiK-net and MeSO-net dataset.

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<sup>233</sup> the residuals by dividing it by the  $\hat{\Sigma}$ , this standard deviation of 0.86 corresponds to an

<sup>234</sup> overestimation of the uncertainty in predictions.

<sup>235</sup> We calculate the  $\hat{\Delta}$  distribution between all observations from both training and <sup>236</sup> validation events with respect to the partial ensemble estimators discussed above. All the <sup>237</sup> four combinations are analysed with respect to the standard Gaussian distribution as shown <sup>238</sup> in Figure 5 and 6.



**Figure 6:** Comparison of the distribution of the standardized residuals with the targeted standard Gaussian distribution for the partial ensemble estimators in all four configurations using the KiK-net dataset.



**Figure 7:** Comparison of the distribution of the standardized residuals with the targeted standard Gaussian distribution for the partial ensemble estimators in all four configurations using both the KiK-net and MeSO-net dataset.

The panels (c) and (d) in both figures gives the closest fit to the standard normal distribution and hence, we can understand that the predictive uncertainty at the training

stations are reliable. But while considering the other two panels of both figures, underes-241 timation of the predictive uncertainty is quite significant - as both are much higher than 1 242 (around 40%). However, the overestimation seems to decrease when an additional network 243 is integrated to the KiK-net dataset. The standard deviation of the validation stations on 244 validation events decreased from 1.39 to 1.35 and the validation stations on training events 245 moved closer to 1 when MeSO-net is added to the U-Net training. These results indicate that 246 the increase in spatial density may contribute to a slight decrease in the underestimation of 247 the predicted standard deviation in the interpolated sites. 248

# 249 5.2 Site amplification

We also analysed the effect of a spatially dense seismological network in the site amplification learned by the U-net architecture. The site specific effects are calculated by training another GMM without using the predictive parameters  $x_S$ ,  $y_s$  and  $z_{bedrock}$ . We then approximate the site amplification  $A\hat{m}ps$  learned by subtracting the GMM predictions with and without the site specific predictive parameters. This will provide us with the site amplification maps shown in Figure 7.

The maps are plotted for the models trained with only KiK-net and also trained with both KiK-net and MeSO-net. Both the maps behave very similar to each other and there is no significant effect on adding more station data. This may again lead to the fact that both spatially and temporally dense networks are required in understanding the detail and reliable effects.

When the map (only using KiK-net) is compared with the amplification maps at SA 1s (Figure 8), significant deamplification can be observed. This may be because of the use of  $z_{bedrock}$  as the predictive parameter for site and hence, this may indicate that the sediments in the basin are damping high frequencies. More studies are required to understand the difference in site amplification in different frequencies.



**Figure 8:** Site amplication estimation of PGA in the Kanto basin approximated from averaged mean predictions. a) KiK-net b) KiK-net + MeSO-net. The black triangles indicate the station locations

# <sup>266</sup> 6 Conclusion and Future work

The importance of spatial density in predicting ground motion IMs using the method 267 developed recently by Lilienkamp et al. (2022) for several applications, including probabilistic 268 seismic hazard analysis, has been discussed in this deliverable. We here assessed and discussed 269 the difference in the uncertainty of U-Net's interpolation of IMs between station locations 270 when data from multiple networks rather than a single network is used. The operation of 271 the U-net architecture is studied, and the U-net is trained using only the KiK-net dataset as 272 well as both the KiK-net and MeSO-net databases. The KiK-net stations have an average 273 inter-station distance of 94 km and 65 stations are located within our study region, ie the 274 Kanto basin. The MeSo-net network is much denser than the KiK-net, with an average 275 inter-station distance of 2-3 km and 300 observation stations located entirely within the 276 Kanto basin, hence the difference of considering a high dense network in training can be 277 understood. 278



**Figure 9:** Site amplication estimation of PSA at 1s in the Kanto basin approximated from averaged mean predictions using the KiK-net data. The black triangles indicate the station locations

The U-net is trained using a single network and by integrating two networks within the 279 same region while the validation stations remain the same. The partial ensemble predictions 280 are analysed to understand the difference of prediction with and without including the MeSO-281 The rmse values between the observations and predictions of the partial ensemble net. 282 estimators decreased after considering the dense network for training. The highest decrease 283 in the rmse is found on the category of validation stations on validation events, which 284 shows the effect of the new integrated network on the interpolation of U-Net, however the 285 significance is still questionable as it is within the standard deviation. More dense networks 286 are required to understand the significance of the decrease in the rmse values. We then 287 compared the standardized residual distribution to analyse the accuracy of the predicted 288 standard deviation. If the standard deviation is less than 1, it corresponds to overestimation 289 of the uncertainty and vice-versa. Both the results show an overestimation of uncertainty by 290

<sup>291</sup> 14-15 %, however when we analysed the partial ensemble estimators the uncertainty range <sup>292</sup> seems to decrease slightly and move towards one when both the datasets are used instead <sup>293</sup> of only KiK-net. Although we integrated a spatially dense network along with the KiK-<sup>294</sup> net dataset to compare the results, more networks with data quantity in both spatial and <sup>295</sup> temporal scale is required to analyse and validate our results.

According to the results provided in this report, the error and the uncertainty tends 296 to reduce with the integration of a spatially dense network, however more studies on the 297 changes of IMs with spatial density is required. The site amplification in different frequen-298 cies is needed to be analysed and investigated using different seismological networks. The 299 availability of the K-net data in the Kanto basin adds an advantage of another network 300 availability within the same region. Integration of K-net along with the KiK-net and the 301 MeSO-net will definitely improve the clarity of the results given in this report. Also, we have 302 analysed the PGA values here, but it will be also more interesting and useful to examine the 303 effect of integration of stations in higher periods PSA (T = 1s) because of its probabilistic 304 seismic hazard analysis applications. 305

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